

## Derivation of the Relaxation Method Algorithm

I. Diffusion/reaction of a single species (e.g., O<sub>2</sub>)

The sediment is divided into vertical layers, or boxes. Within a given box, steady state is defined as the condition where the sum of the fluxes sum to zero; the important fluxes are diffusion and reaction. We define a few units of currency:

$$inv \left[ \frac{\text{mol}}{\text{cm}^2} \right] = conc \left[ \frac{\text{mol}}{\text{l pw}} \right] \cdot \Delta z \cdot \Phi \cdot 10^{-3} \quad (\text{A1.b.1})$$

where *inv* is the inventory of a solute within a box, per unit area on the sediment surface, and *conc*, its pore-water concentration.

**Derivation of the diffusion expression.** Using the currency of the solute inventory, the diffusive and reaction fluxes into and out of the box at level "i" are,

$$\begin{aligned} 0 = & D \frac{F_{i-1} + F_i}{2} \frac{2}{\Delta z_{i-1} + \Delta z_i} (conc_{i-1} - conc_i) \cdot 10^{-3} \\ & - D \frac{F_i + F_{i+1}}{2} \frac{2}{\Delta z_i + \Delta z_{i+1}} (conc_i - conc_{i+1}) \cdot 10^{-3} \\ & + J \left[ \frac{\text{mol}}{\text{l tot s}} \right] \cdot 10^{-3} \cdot \Delta z \end{aligned} \quad (\text{A1.b.2})$$

where the subscript "i-1" indicates the box above and i+1 the box below. F is the formation factor (see chapter 2), Δz is the thickness of the box [cm], and J is the reaction rate, in moles/liter total/second. When the sum of these fluxes for each box i equals zero, as above, then the system is in steady state.

The expression converted to the currency of concentrations [moles/l pore water]:

$$\begin{aligned}
0 = & D \frac{F_{i-1} + F_i}{2} \frac{2}{(\Delta z_{i-1} + \Delta z_i) \Delta z_i} \frac{1}{\Phi} (\text{conc}_{i-1} - \text{conc}_i) \\
& - D \frac{F_i + F_{i+1}}{2} \frac{2}{(\Delta z_i + \Delta z_{i+1}) \Delta z_i} \frac{1}{\Phi} (\text{conc}_i - \text{conc}_{i+1}) \\
& + \frac{j}{\Phi}
\end{aligned} \tag{A1.b.3}$$

which is obtained by substituting (A1.b.1) into (A1.b.2).

To simplify the notation, I define

$$D_- = D \frac{F_{i-1} + F_i}{2} \frac{2}{(\Delta z_{i-1} + \Delta z_i) \Delta z_i} \frac{1}{\Phi} \tag{A1.b.4}$$

$$D_+ = D \frac{F_i + F_{i+1}}{2} \frac{2}{(\Delta z_i + \Delta z_{i+1}) \Delta z_i} \frac{1}{\Phi} \tag{A1.b.5}$$

$$x_i = \text{conc}_i \tag{A1.b.6}$$

Then equation (A1.b.3) becomes,

$$0 = D_- (x_{i-1} - x_i) - D_+ (x_i - x_{i+1}) + \frac{j}{\Phi} \tag{A1.b.7}$$

**Iteration to a solution.** Since the first guess for the steady state concentrations,  $x_i$ , to the diffusion/reaction equation is usually incorrect, we define

$$R_i = D_- (x_{i-1} - x_i) - D_+ (x_i - x_{i+1}) + \frac{j}{\Phi} \neq 0 \tag{A1.b.8}$$

where the residual term  $R_i$  is defined for all boxes,  $i$ .

The solution is found by iteration, using a procedure analogous to Newton's method for finding the solution to

$$f(x) = 0 \quad (\text{A1.b.9})$$

on an interval. Each successive guess ( $x + \Delta x$ ) for the value of  $x$  to satisfy this expression is determined using the last guess ( $x$ ), by finding the  $y$  intercept of the tangent to the curve  $f(x)$ , so that  $\Delta x$  satisfies

$$\left. \frac{\partial f}{\partial x} \right|_{x_1} \cdot \Delta x + f(x_1) = 0 \quad (\text{A1.b.10})$$

For the system of solute concentrations in a vertical stack of  $n$  boxes in the sediment, there is a system of  $n$  equations with  $n$  unknown  $\Delta x_i$  values. Each concentration ( $x_i$ ) appears in three equations, the ones describing the fluxes for the box itself, plus above and below ( $\Delta x_i, \Delta x_{i-1}, \Delta x_{i+1}$ ). A generalized equation is

$$\frac{\partial R_i}{\partial x_{i-1}} \Delta x_{i-1} + \frac{\partial R_i}{\partial x_i} \Delta x_i + \frac{\partial R_i}{\partial x_{i+1}} \Delta x_{i+1} = -R_i \quad (\text{A1.b.11})$$

**Boundary conditions.** At the sediment surface, the solute concentration is fixed at the bottom water value

$$x_1 = x_{bw} \quad (\text{A1.b.12})$$

During the iteration, the concentration at  $i=1$  remains constant, so that,

$$\Delta x_1 = 0. \quad (\text{A1.b.13})$$

Accordingly, the expression for  $i=2$  analogous to expression (A1.b.11) is,

$$\frac{\partial R_2}{\partial x_2} \Delta x_2 + \frac{\partial R_2}{\partial x_3} \Delta x_3 = -R_2. \quad (\text{A1.b.14})$$

At the bottom boundary ( $i = n$ ), the imposed condition is that the diffusive flux equals zero; that is

$$x_{n+1} = x_n \quad (\text{A1.b.15})$$

and, during the iteration,

$$\Delta x_{n+1} = \Delta x_n. \quad (\text{A1.b.16})$$

Therefore, the constraining equation from the bottom box ( $i = n$ ) is,

$$\frac{\partial R_n}{\partial x_{n-1}} \Delta x_{n-1} + \left( \frac{\partial R_n}{\partial x_n} + \frac{\partial R_n}{\partial x_{n+1}} \right) \Delta x_n = -R_n. \quad (\text{A1.b.17})$$

The  $n$  equations for the  $n$  unknowns  $\Delta x_{(1 \rightarrow n)}$  are solved simultaneously using Gauss-Jordan elimination (Press et al., 1986). For a simple system consisting of only 4 depth levels, the matrices are constructed as follows:

$$\begin{array}{cccc|ccc}
 \frac{\partial R_2}{\partial x_2} & \frac{\partial R_2}{\partial x_3} & 0 & 0 & \Delta x_2 & & -R_2 \\
 \frac{\partial R_3}{\partial x_2} & \frac{\partial R_3}{\partial x_3} & \frac{\partial R_3}{\partial x_4} & 0 & \Delta x_3 & & -R_3 \\
 0 & \frac{\partial R_4}{\partial x_3} & \frac{\partial R_4}{\partial x_4} & \frac{\partial R_4}{\partial x_5} & \Delta x_4 & = & -R_4 \\
 0 & 0 & \frac{\partial R_5}{\partial x_4} & \frac{\partial R_5}{\partial x_5} + \frac{\partial R_5}{\partial x_6} & \Delta x_5 & & -R_5
 \end{array}$$

(A1.b.18)

## II. pH-Coupled Carbonate System

In contrast to the equation describing the distribution of  $O_2$  in sediment pore water (described above) the equations describing  $CO_2$ ,  $HCO_3^-$ , and  $CO_3^{=}$  are coupled together by the carbonate buffer system. Three constraints are used to define the state of the system within each box: the  $\Sigma CO_2$  balance, the alkalinity balance, and the requirement of pH equilibrium. Thus, there are three equations for each box, and three unknowns, representing the  $\Delta x$ 's for the three species.

As for the treatment of  $O_2$ , above, subscripts following a variable indicate the box number or depth level. Superscripts following a variable denote the species, as,

$$x^1 = CO_2$$

$$x^2 = HCO_3^-$$

$$x^3 = CO_3^{=}$$

Superscripts leading a variable indicate that the variable is associated with one of the three constraints described above, so that

$${}^1R_i,$$

for example, is the residual of the  $\Sigma CO_2$  equation, described below.

**k = 1:  $\Sigma\text{CO}_2$  balance.** In steady state, the diffusion of all three species should equal the sum of the inputs of carbon by organic carbon oxidation and calcite dissolution:

$${}^1R_i = \sum_{j=1}^3 D_-^j (x_{i-1}^j - x_i^j) - D_+^j (x_i^j - x_{i+1}^j) + \frac{J^1}{\Phi} + \frac{J^3}{\Phi} \quad (\text{A1.b.19})$$

**k = 2: alkalinity balance.** The alkalinity balance in the code is normalized to units of calcite dissolution, so that the quantity conserved is twice the alkalinity:

$${}^2R_i = \frac{1}{2} \left[ D_-^2 (x_{i-1}^2 - x_i^2) - D_+^2 (x_i^2 - x_{i+1}^2) \right] + D_-^3 (x_{i-1}^3 - x_i^3) - D_+^3 (x_i^3 - x_{i+1}^3) + \frac{j^3}{\Phi} \quad (\text{A1.b.20})$$

**k = 3: pH equilibrium.** The three species are required to be in pH balance in each box:

$${}^3R_i = \frac{x_i^1 \cdot x_i^3}{(x_i^2)^2} - \frac{K_1}{K_2} \quad (\text{A1.b.21})$$

where the denominator in the first term is squared (an abundance of superscripts!).

**Solution by iteration.** The system of equations is solved by iteration similarly to the case of a single species, above. The general equation is:

$${}^kR_i = \sum_{j=1}^3 \left( \frac{\partial {}^kR_i}{\partial x_{i-1}^j} \Delta x_{i-1}^j + \frac{\partial {}^kR_i}{\partial x_i^j} \Delta x_i^j + \frac{\partial {}^kR_i}{\partial x_{i+1}^j} \Delta x_{i+1}^j \right). \quad (\text{A1.b.22})$$

When the summation term is expanded, there are nine terms, constraining nine unknown  $\Delta x$ 's.

In order to simplify the notation, the following abbreviation will be used in the text and in the code.

$$\text{define: } \frac{\partial^k R_i}{\partial x_1^j} \equiv \text{DR}(K,J,L,I) \quad (\text{A1.b.24})$$

k = equation number:                    1 =  $\Sigma\text{CO}_2$   
     2 = alkalinity  
     3 = pH equilibrium

j = species number:                    1 =  $\text{CO}_2$   
     2 =  $\text{HCO}_3^-$   
     3 =  $\text{CO}_3^{=}$

l = level offset:                        1  $\rightarrow$  i-1 (level above)  
     2  $\rightarrow$  i  
     3  $\rightarrow$  i+1 (level below)

i = depth level

Using this notation, for example, the FORTRAN term for the derivative of the residual of the  $\Sigma\text{CO}_2$  equation at depth level number 7, with respect to changes in the bicarbonate concentration in box number 7, is

$$\frac{\partial \Sigma\text{CO}_2 R_7}{\partial \text{HCO}_3 (7)} \equiv \text{DR}(1,2,2,7)$$

A sample of the matrix construction for the simultaneous solution of these equations can be presented by further shortening the notation to:

$$1227 \equiv DR(1,2,2,7).$$

The matrix appropriate to a system four boxes deep (including  $i=1$  for bottom water) would be given by:

1122	1222	1322	1132	1232	1332	0	0	0	$\Delta x_2^1$	${}^1R_2$
2122	2222	2322	2132	2232	2332	0	0	0	$\Delta x_2^2$	${}^2R_2$
3122	3222	3322	3132	3232	3332	0	0	0	$\Delta x_2^3$	${}^3R_2$
1113	1213	1313	1123	1223	1323	1133	1233	1333	$\Delta x_3^1$	${}^1R_3$
2113	2213	2313	2123	2223	2323	2133	2233	2333	$\Delta x_3^2$	${}^2R_3$
3113	3213	3313	3123	3223	3323	3133	3233	3333	$\Delta x_3^3$	${}^3R_3$
0	0	0	1114	1214	1314	1124	1224	1324	$\Delta x_4^1$	${}^1R_4$
						+1134	+1234	+1334		
0	0	0	2114	2214	2314	2124	2224	2324	$\Delta x_4^2$	${}^2R_4$
						+2134	+2234	+2334		
0	0	0	3114	3214	3314	3124	3124	3324	$\Delta x_4^3$	${}^3R_4$
						+3134	+3234	+3334		

(A1.b.25)