STABILITY OF HIBONITE AND CaAl₄0₇ IN THE SOLAR NEBULA; L. Grossman¹, C.A. Geiger¹, O.J. Kleppa¹, B.O. Mysen², and J.M. Lattimer³. Dept. of the Geophysical Sciences, University of Chicago, Chicago, IL 60637. Geophysical Laboratory, Carnegie Institution of Washington, Washington, DC 20008. Dept. of Earth & Space Sciences, SUNY, Stony Brook, NY 11794.

Many calcium aluminates are known, but only two are found in meteorites. Hibonite, CaAl₁₂O₁₉, has been found in many refractory inclusions in C2 and C3 chondrites (1,2). CaAl₄O₇, in contrast, has been reported from only 3 refractory inclusions (3,4,5). Equilibrium thermodynamic calculations of the sequence of condensation of phases from a cooling gas of solar composition did not consider calcium aluminates until after free energy data for these compounds were reported by Allibert et al. (6). Fegley (7), Davis et al. (8) and Kornacki & Fegley (9) computed that, at 10⁻³ atm total pressure, corundum condenses first at ∿1740 K and reacts with gaseous Ca to form hibonite 10-20° lower in temperature. Calculations in (7) and (9) show that, at a still lower temperature, hibonite reacts with the gas to form CaAl₄O₇. Kornacki & Fegley (9), as a result, used the absence of CaAl₄O₇ from hibonite-bearing inclusions as an argument that the latter are not high-temperature condensates but, instead, formed by partial melting and distillation of dust aggregates. MacPherson & Grossman (10) pointed out, however, that CaAl₄O₇ would disappear completely from the condensation sequence if its free energy were only 6.1 kJ/mole less stable than the value in (6) and suggested that the absence of CaAl₄O₇ from condensate assemblages may be due to incorrect thermodynamic data for this phase. Since differences larger than 6.1 kJ/mole exist in the reported data for CaAl₄O₇ and since no direct calorimetric data exist for hibonite, we determined the enthalpies of formation of CaAl₄O₇ and hibonite by oxide melt solution calorimetry.

Synthetic CaAl₄O₇ and CaAl₁₂O₁₉ were prepared by heating intimate, stoichiometric mixtures $CaAl_{12}O_{19}$, has been found in many refractory inclusions in C2 and C3 chondrites (1,2).

Synthetic $CaA1_40_7$ and $CaA1_{12}0_{19}$ were prepared by heating intimate, stoichiometric mixtures of $CaC0_3$ and $A1_20_3$ at $1250^{\circ}C$ overnight. These mixtures were then subjected to a number of cycles of high-temperature sintering and intermediate grinding operations. This is the same synthesis technique used to prepare these compounds in all previous studies of their thermodynamic propertechnique used to prepare these compounds in all previous studies of their thermodynamic properties (6,11,12,13), except that care was taken here to use more intermediate grinding steps, far longer total heating times, far longer heating times above 1500°C and usually higher maximum temperatures than in all previous work. For CaAl $_4$ 07, 11 cycles were used, ranging from 1 cycle of 240 hr at 1400°C to 5 cycles of 4-6 hr each at 1650°C. For CaAl $_1$ 20 $_1$ 9, 9 cycles between 1500 and 1625°C were used, totalling 400 hr of heating. In all previous thermodynamic studies, synthetic calcium aluminates were found to be pure by X-ray powder diffraction. X-ray study showed no impurities in our synthetic CaAl $_4$ 07 but minor amounts of α -Al $_2$ 03 in the hibonite. SEM study of our powders, however, revealed hibonite and a calcium silicate in the synthetic CaAl $_4$ 07, and corundum in the synthetic hibonite. Excess Al $_2$ 03 in the hibonite sample and SiO $_2$ in the CaAl $_4$ 07 were probably introduced during grinding in Al $_2$ 03 and agate mortars, respectively. SEM study of CaAl $_4$ 07 run products thus revealed significant impurities undetected by X-ray diffraction, suggesting that the latter technique may have been insufficient to establish the purity of calcium aluminates prepared for earlier thermodynamic studies (6,11,12,13). From point-count modes of SEM photos, we determined that the heat of solution value obtained for the hibonite sample had to be corrected for the presence of 8.5 wt.% corundum and the value for the CaAl $_4$ 07 sample had to be corrected for the presence of 8.5 wt.% corundum and the value for the CaAl $_4$ 07 sample had to be corrected for the presence of 8.5 wt.% corundum and the value for the CaAl $_4$ 07 sample had to be corrected for the presence of 8.5 wt.% corundum and the value for the CaAl $_4$ 07 sample had to be corrected for 8.9 wt.% hibonite and 2.1% silicate, assumed to be wollastonite.

Using a Ni-block, Calvet-type, twin calorimeter (14,15), 40 mg samples of powdered CaO, Al $_2$ 03, CaAl $_4$ 0

digital integration of the temperature-time curve generated by two 48-couple Pt-Pt13Rh thermopiles. Frequent calibration was done by dropping small pieces of gold into the calorimeter and measuring the heat effect as the gold rose from room temperature to calorimeter temperature. After correction for impurities, the heat of formation of each of the calcium aluminates from its oxides at 1063 K, Δ H 2 (1063), was calculated by subtracting its heat of solution from those of its constituent oxides in stoichiometric proportion. Δ H 2 (1063) = -25.6 \pm 4.7 kJ/gfw for CaAl $_4$ O $_7$ and

constituent oxides in stoichiometric proportion. ΔH²(1063) = -25.6 ± 4.7 kJ/gfw for CaAl₄07 and -33.0 ± 9.7 kJ/gfw for CaAl₁2019.

Kornacki & Fegley (9) used free energy data for CaAl₄07 and CaAl₁2019 from the electrochemical study of Allibert et al. (6) in their condensation calculations. Those data were obtained in the temperature range 900-1200 K and had to be extrapolated to 1600-1750 K, the range of condensation temperatures. Gibbs free energy data for these 2 phases have since become available from the later electrochemical study of Kumar & Kay (13) over the temperature range 1100-1500 K. These data are preferred over those in (6) as the uncertainties are a factor of 2 smaller and the extrapolation to condensation temperatures 250° less than in those of (6). Using the equations given in (13) for the temperature variation of the free energies of formation from the oxides, ΔG¾ = -69.00 ± 0.84 kJ/gfw for hibonite and -47.67 ± 1.26 kJ/gfw for CaAl₄07 at 1300 K, the midpoint of the temperature range investigated experimentally in (13).

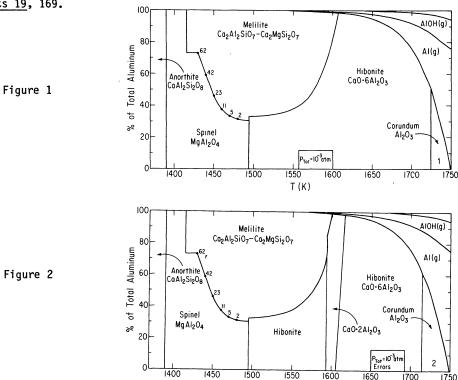
Using heat capacity data from (17) and (18), we make the small temperature correction to our measured ΔH¾(1063) to yield ΔH¾(1300) = -28.78 ± 5.15 kJ/gfw for CaAl₄07 from the oxides. Using this value and the above ΔG¾(1300) from (13), we calculate ΔS¾(1300) = [ΔH¾(1300) - ΔG¾(1300)]/1300 = 14.5 J/K gfw. Assuming ΔG¾(1300) which differ by more than a factor of 2 from the intercept and slope, respectively, of the ΔG¾ vs. T equation for CaAl₄07 in (13). Thus, the equation in (13) cannot be used for temperature extrapolation. We calculate ΔG¾(1600) = -52.03

equation in (13) cannot be used for temperature extrapolation. We calculate $\Delta G_{\rm f}^2(1600)$ = -52.03 \pm 1.95 kJ/gfw for CaAl₄O₇ from $\Delta G_{\rm f}^2(1300)$ from (13) and our calculated value of $\Delta S_{\rm f}^2(1300)$. Without heat capacity data for hibonite, we assume $\Delta H_{\rm f}^2(1300) \sim \Delta H_{\rm f}^2(1063)$ determined experi-

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mentally above and calculate $\Delta S_f^2(1300) = 27.7$ J/K gfw as above. Again, $\Delta H_f^2(1300)$ and $\Delta S_f^2(1300)$ determined here are greatly different from the intercept and slope of the equation for hibonite in (13). For hibonite, we calculate $\Delta G_f^2(1600) = -77.31 \pm 2.46$ kJ/gfw from $\Delta G_f^2(1300)$ from (13) and our value for $\Delta S_f^2(1300)$.

Using the same technique as in Lattimer & Grossman (19) but with an updated set of thermodynamic data including free energies for hibonite and CaAl₄O₇ determined herein, we cal₂ulated the sequence of condensation of phases from a cooling gas of solar composition. At 10⁻³ atm total pressure, corundum condenses first at 1749 K and disappears at 1725 K where it reacts with gaseous Ca to form hibonite, as shown in Fig. 1. Perovskite condenses at 1675 K. Hibonite begins to react with gaseous Ca and Si to form gehlenite at 1607 K. Hibonite in excess of this reaction reacts totally to form spinel at 1494 K. The sequence of reactions is the same at 10⁻⁴ and 10⁻³ atm total pressure. Thus, use of the new thermodynamic data yields no stability field for CaAl₄O₇ in a solar gas over the pressure range 10⁻³-10⁻⁵ atm. Fig. 2 shows the maximum possible stability field of CaAl₄O₇ in a solar gas at 10⁻³ atm, obtained by making CaAl₄O₇ as stable and hibonite as unstable as the error bars on their free energies permit. In this case, hibonite first appears by reaction of corundum with gaseous Ca at 1715 K. Hibonite begins to react with gaseous Ca to form CaAl₄O₇ at 1616 K and disappears at 1605 K. CaAl₄O₇ begins to react with gaseous Ca and Si to form melilite at 1600 K and the CaAl₄O₇ in excess of this reaction is completely converted back to hibonite at 1593 K. After this, hibonite gradually reacts with the gas to form gehlenite, finally disappearing at 1495 K. Thus, even in this extreme case, the stability field of CaAl₄O₇ is only 23° wide. It is unlikely that CaAl₄O₇ has a stability field in a gas of solar composition. Certainly, its absence from refractory inclusions may be the result of metastable condensation. As the stability field of CaAl₄O₇ in refractory inclusions may be the result of metastable condensation. As the stability field of CaAl₄O₇ in refractory inclusions may be the result of metastable condensation. As the stability field of CaAl₄O₇ in refractory



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